EFFECT OF PHYSICAL AND CHEMICAL PROPERTIES OF NORMAL SOILS ON HYDRAULIC CONDUCTIVITY VALUES AS MEASURED BY DIFFERENT METHODS

Saffan, M.M; M.R. Khalifa; A.M. Abou El-Khir and I.A. El-Saiad

Soil Sci. Dept. Fac. of Agric. Kafr El-Sheikh, Tanta Univ., Egypt

ABSTRACT

Four soil profiles representing normal soils(non-saline non-sodic) were taken at Byalla district (Kafr El-Sheikh governorate) to study the relationship between physical and chemical properties of the soil and values of hydraulic conductivity as determined by three different methods (constant head, falling head and auger hole). Simple and multiple linear regression relationships were conducted between values of saturated hydraulic conductivity (Ks) and the different physical and chemical properties of the soil .The important results could be summarized as follows:

- 1- Ks as measured by constant head was affected by aggregation conditions and showed a highly significant positive correlation with each of water stable aggregates >0.25 mm ($R^2 = 0.6814^{**}$), aggregation index ($R^2 = 0.6383^{**}$) and mean weight diameter ($R^2 = 0.6439^{**}$). A highly significant positive correlation was found between Ks and both soluble Ca⁺⁺($R^2 = 0.6309^{**}$) and exchangeable Mg ($R^2 = 0.5577^{**}$). Multiple linear regression between the above- mentioned chemical properties of the soil and Ks showed a significant correlation coefficient ($R^2=0.9340^{*}$). It is obvious that soluble Ca⁺⁺ is the most effective chemical parameter which affect Ks.
- 2- Ks determined by falling head method was greatly affected by aggregation parameters and has a highly positive correlation with both water stable aggregates >0.25 mm ($R^2=0.6384^{**}$) and structure coefficient SC ($R^2=0.7409^{**}$). A highly negative correlation was found between water conductivity and soil bulk density ($R^2 = 0.7671^{**}$). The multiple linear regression between the abovementioned physical properties and Ks revealed a significant correlation ($R^2=0.9250^{*}$). Ks by falling head method was also affected by the chemical properties of the soil. A highly positive correlation was found between values of hydraulic conductivity and each of exchangeable Mg ($R^2=0.9114^{**}$), cation exchangeable capacity ($R^2=0.7074^{**}$), calcium carbonate content ($R^2=0.8106^{**}$) and organic mater content ($R^2=0.999^{*}$). Multiple linear regression between (Ks) and the chemical properties of the investigated soils showed a significant correlation coefficient ($R^2=0.999^{*}$).

- 3- Statistical analysis showed no significant correlation between the values of saturated hydraulic conductivity which obtained by both constant head and falling head methods, and by the auger-hole method .
- Key words: Hydraulic conductivity, constant head, falling head, auger hole, normal soils, non-saline non-sodic soils, physical and chemical properties

INTRODUCTION

Hydraulic conductivity (Ks) is a common and essential parameter for most of soil physical investigations due to the connected relation of (Ks) to many of soil water relations as water conservation, irrigation scheduling, drainage planning and soil-plant relationships . Since many methods for determining (Ks) can be conducted in the field as well as in the laboratory, the obtained data could be varied for the same soil. Usually, measurements of hydraulic conductivity in the lab. or in the field is tedious, expensive and require excessively large numbers of readings. Many scientists and researchers carried out many investigations about soil water conductivity as affected by different physical and chemical properties of the soil (Camplell and Camplell, 1982: Saffan, 1984; Saxton et al., 1986; Gallichand et al., 1990; El-Samanoudi, 1992; Yosry, 1992;Abdel-Nasser, 1995; Chen et al., 1998; Arya et al., 1999 and El-Henawy, 2000).

Researchers may have developed or used one method for measuring water hydraulic conductivity on definite soil samples. Therefore, the obtained data for (Ks) may be exchanged as another method for the same sample was used. Therefore, the main objectives of this study were to :

- 1- Measure saturated hydraulic conductivity for normal soils(non-saline non-sodic) with three different methods (constant head, falling head and auger-hole methods).
- 2- Calculate the relationship and correlation between the different obtained values for a definite soil sample and its physical and chemical properties.
- 3- Establish correlation coefficient that describe the relationships between (Ks) values which obtained by constant head, falling head and auger-hole methods.

MATERIAL AND METHODS

Four soil profiles from Kafr El-Sheikh governorate (Byalla district)
were chosen in different locations representing normal soils (non-saline non-sodic soils according to Richareds, 1954 and James et al., 1982).

Disturbed and undisturbed soil samples were taken from the successive soil depths of 0-30; 30-60 and 60-90 cm, respectively. Undisturbed soil samples were taken using core samples to determine bulk density, hydraulic conductivity and aggregate size distribution. Disturbed

soil samples were air dried, gently ground with wood paste, sieved through 2 mm sieves and kept in plastic bags for soil analysis. Physical and chemical properties were determined using the following methods :

Particle size distribution was carried out according to Gee and Bauder (1986). Aggregate size distribution was determined by wet sieving technique(Yoder, 1936) using sieves with 2.00, 1.00, 0.50, 0.25 mm screen opening according to klute (1986). Total of water stable aggregation (Total of W.S.A) was calculated according to Ibrahim, (1964); aggregation index (AI), and mean weight diameter (MWD) according to Baver et al., (1972). Optimum size of aggregates (Opt.size) was measured after Kohnke, (1968) and structure coefficient (SC) according to El-shafei and Ragab, (1975) Bulk density (B_d) was determined using the method described by Vomocil, (1957). The saturated hydraulic conductivity was determined by three methods: constant head method (Klute, 1986), auger hole method to a depth of 90 cm (Van Beers, 1958) and falling head method (Klute, 1986). ECe of the soil saturation extract, soluble cations and anions, pH, organic matter, cation exchange capacity, exchangeable cations and total calcium carbonate were determined according to the conventional methods described by (Jackson, 1958 and piper, 1950).

Correlation coefficients and multiple regression between saturated hydraulic conductivity and each of the studied properties were done after Snedecor and Cochran,(1967).

Results and Discussion

1- General soil characteristics :

The values of some physical and chemical properties are shown in Tables 1 and 2 respectively Data revealed that, the bulk density (B_d) ranged between 1.12 to 1.27 and 1.29 to 1.36gm/cm³ reveal in (0-30) and (30-60) cm respectively. Generally B_d values increased with depth in all soil profiles. Total carbonate content of all the soil samples in different profiles , indicated no calcification as it ranged between 0.9 to 2.30% and, decreased with increasing soil depth in all profiles studied. Data in Table 1 showed that , soil texture in all of the profiles is clayey , whereas the clay, silt and sand content ranged between 51.4 to 52.2%, 25.3 to 25.9% and 20.3 to 24.9% in 0-30 cm respectively.

Data in Table (1) showed that total of water stable aggregates > 0.25 mm was ranged between 70.04 to 75.72%,(in 0-30 cm). In addition its values decreased with increasing soil depth. Values of aggregation index (A.I), mean weight diameter (MWD) and structure coefficient (SC) in (0-30 cm) ranged between 0.561 to 0.840, 1.121 to 1.668 mm⁻ and 2.34 to 3.12. respectively.

	Profiles						files	les						
Soil Properties				۲		1	P ₂			P,			P.	
			Depth , Ст											
			0-30	30-60	60-90	0-30	30-60	60-90	0-30	30-61)	60-90	0-30	30-60	60-90
	Constant-head method cm h ⁻¹		0.24	0.24	0.18	0.24	0.18	0.12	0,24	0.24	0.12	0.30	0,24	0.18
Ks	Falling-bead method cm h ⁻¹ Auger-hole method cm h ⁻¹		0.18	0.15	0.12	0.21	0.16	0.12	0.19	0.16	0.12	0.22	0,15	0.13
				0.07			0.10			0.08			0.10	.
Bulk	Bulk density gm/cm ¹		1.27	1.32	1,39	1.12	1.29	1.31	1.20	1.36	1.35	1.17	1.29	1.32
	CaCo ₃ %		2.20	1.20	0.80	1.80	1.60	1.00	2.00	1.40	0.90	2.30	1.10	0.60
		Send %	22.70	23.40	22.90	22.50	22.60	22.40	22.60	23.00	22.40	22.60	23.40	22.60
Practical size*	distribution	Silt %	25.90	27.30	26.10	25.30	24.70	25.50	25.60	26.00	25.00	25.70	27.10	25.90
	Clay		51.40	49.30	51.00	52.20	52.70	54.10	51.80	51.00	52.60	51.70	49.50	51.50
Total of W.S.A %		70.04	68.08	56.31	75.72	71.82	50.83	72.88	68.58	57.75	72.82	69.08	56.35	
aggregation index (AI)			0.56	0.69	0.30	Ü.84	0.83	0.28	0.70	0.76	0.30	0.84	0.69	0.30
MWD m.m			1.121	1.381	0.597	1.668	1.654	0.560	1.396	1.513	0.592	1.680	1.374	0.596
Optimum size of aggregates (OP.size)			56.79	51.34	55.77	50.38	49.47	47.88	53.59	50.41	51.83	49,47	51.34	55.77
structure coefficient(SC)			2.34	2.13	1.29	3.12	2.55	1.03	2.73	2.18	1.37	2.52	2.10	1.27

Table 1. Saturated hydraulic conductivity (Ks) as measured by different methods and some physical properties of the studied soil profiles.

* The soil texture is clayey in all the studied samples.

·····	Profiles									:			
Soil Properties		Pi			P2			P3			P4		
		Depth , Cm											
		0-30	30-60	60-90	0-30	30-60	60-90	0-30	30-60	60-90	0-30	30-60	60-90
PH (1:2.5) soil : water		8.28	8.17	8.27	8.12	8.22	8.19	8.17	8.20	8.19	8.18	8.21	8.13
EC _c dS/m		2.31	3.27	3.99	2.84	2.48	3.33	2.58	2.88	3.68	2.35	3.25	3.98
	Na	2.33	1.03	4.73	7.00	7.18	9.42	0.43	3.52	3.93	2.82	3.30	2.12
Soluble cations	Ca ⁺⁺	13.00	17.60	7.40	13.90	10.05	10,30	13.90	13.90	9.30	14.80	15.50	5.60
meq/L	Mg ⁺⁺	7.40	13.90	27.80	5.60	7.30	13.20	11,10	11.10	23.20	5.60	11.90	32.00
	K	0,37	0.37	0.37	1,90	0.27	0.38	0,37	0.28	0.37	0.28	1.80	0.38
	Co 3	1.90	-	0.93	1.90	1.80	1.90	1,90	0.93	0.93	1.90	1.80	0.94
Soluble anions	HCO3	16.7	9.38	12.1	7.40	5.50	9.40	14,90	10.20	12.10	9.30	7.30	5.60
meq/L	Cl	0.46	15.00	15.60	13,00	10.10	13.00	4.60	10.30	15:70	10.20	12.30	17.80
·	SO'	4.04	8.10	11.70	6,10	7.40	9.00	4.40	7.40	8.10	2.10	11.10	15.66
0.M %		1.68	0.65	0.56	1,90	0.63	0.53	1.80	0.64	0.55	1.50	0.63	0.54
Enstrumentale	Na	2.30	3.30	3.50	1.50	2.20	2.60	2.40	2.80	3.30	2.60	2.70	3.40
Exchangeable	Ca ⁺⁺	20,50	15.90	15.70	22.60	13.90	17.10	21.60	14.90	16.40	20.30	15.60	15.50
cations meq/100 gm soit	Mg ⁺⁺	13.00	11.10	9,90	15,80	9,90	8.90	14.40	10.50	9.40	16.00	11.00	9,70
50H	K ⁺	1.12	1.02	1.16	0.75	0.86	1.27	0,89	0.94	1.22	1.27	1.7	1.24
C.E.C.		37.22	31.42	30.66	41.35	27.66	30.77	39.29	29.54	30.72	41.37	31.40	30.64
Ex. Ca P		55,10	.50.60	51.20	54.70	50.30	55.60	54.90	50.50	53.40	55.00	50.60	55.80
Ex. MgP		34.90	35.30	31.10	38.20	35.80	28.90	36.60	35.60	30.00	34.60	35.00	30.90
ESP		6.20	10.50	11.40	3.60	7.90	8.50	6.10	9.50	10.70	6.30	8.60	11.10
Soit diagnosis		Normal	Normal	Normal	Normal	Normal	Normal	Normal	Normal	Normai	Normal	Normal	Normai

Table 2. Some chemical properties of the studied soils profiles .

Data in Table (1) revealed that Ks values determined by constant head method were ranged between 0.24 to 0.30 cm/h in (0-30 cm), while the corresponding values of Ks determined by falling head method were ranged between 0.18 to 0.22 cm/h. On the other hand Ks values determined by auger hole method were ranged between 0.07 to 0.1 cm/h in the studied soil profiles.

Data in Table (2) showed that electrical conductivity (ECe) values, ranges between 2.31 and 2.84 dS/m, pH values fluctuate between 8.12 and 8.28, organic matter percent, ranges from 1.50 to 1.90%, and the soluble Na⁺ values were lesser than the summation of soluble Ca⁺⁺ and Mg⁺⁺⁽ in 0-30 cm). Also data in Table (2) revealed that in (0-30)cm exchangeable sodium percentage (ESP) was less than 15% and it's values range between (3.6 and 6.1%), Ex. Ca P values ranged between 54.7 and 55.1%, Ex. MgP ranged between 34.6 and 38.2% and C.E.C. fluctuates between 37.22 and 41.37 meq/100 g soil. Exchangeable Ca or Mg predominate on the exchangeable complex in the ail studied profiles.

2- The relation between saturated hydraulic conductivity (Ks) by constant head methods and some soil variables :

Ks has highly positive significant correlation with aggregation index (AI) ($R^2 = 0.6323^{**}$), mean weight diameter (MWD) ($R^2 = 0.6439^{**}$) and total of water stable aggregates > 0.25 mm (WSA %) ($R^2 = 0.8814^{**}$), while it had a positive significant correlation with structure coefficient (S.C) ($R^2 = 0.5524^{*}$) (Table, 3).

Multiple regression equation was calculated between the investigated Ks and the four independent variables of soil structure was: $K_s = -0.414-26.860$ (AI) + 13.41 (MWD)-0.0553 (S.C) + 0.0116 (WSA%) (R² = 0.775, insignificant).

	Particular	Linear regression	\mathbf{R}^2	
Y _	X	equations		
	WSA %	Y = -0.208 + 0.0063 x	0.6814	
	AI	Y = 0.0727 + 0.2211 x	0.6323**	
Ks cm/ h	MWD (mm)	Y = 0.0711 + 0.42 x	0.6439**	
	S.C	Y = 0.0623 + 0.0699 x	0.5524*	
	Exchangeable Mg ⁺⁺ (meq/100 g soil)	Y = 0.0129 + 0.0169 X	0.5877**	
	CaCO ₃ %	Y = 0.092 + 0.08 X	0.4921*	
	Soluble Ca ⁺⁺ (meq/L)	Y = -0.0192 + 0.0178 X	0.6509*	
	Soluble Mg ⁺⁺ (meq/L)	Y = 0.2971 - 0.0074 X	0.4433*	

Table 3. The linear regression between Ks determined by constant head method and some physical and chemical properties .

** Highly significant at level 1%

Significant at level 5%.

The obtained results are in agreement with those found by; Zein El-Abedine (1979), Talha et al. (1979), Abdel-Rasoul et al. (1999). They reported that aggregation parameters of soil are significantly and positively related with the Ks.

Ks had highly positive significant correlation with soluble calcium $(R^2 = 0.6509^{**})$ and a exchangeable magnesium $(R^2 = 0.5877^{**})$ while it had a positive significant correlation with calcium carbonate % $(R^2 = 0.4921^*)$. On the other hand, soluble magnesium had a negative significant correlation with Ks $(R^2 = 0.4433^*)$, (Tables 3). Multipleregression equation between (Ks) and the independent variables of chemical properties, could be expressed as follows :

Ks = -0.110 + 0.014 (soluble Ca⁺⁺)^{*} - 5.77 E-04 (soluble Mg⁺⁺) + 1.581 E-05 (exchangeable Mg⁺⁺) + 0.0518 (CaCO₃ %). (R² = 0.934^{*}).

It's obvious that, soluble Ca⁺⁺ had the most important parameter which can effect the saturated water flow in the soil. The obtained results are confirmed with those found by Mustafa (1969) who pointed out that the effect of different chloride salts on the permeability coefficient could be arranged in the following order: CaCl₂ > MgCl₂ > NaCl. Oster et al. (1980) found that the suitable presence of divalent ions increase the soil hydraulic conductivity. The same conclusion was found by Madkour et al. (1999).

3- The relation between saturated hydraulic conductivity (Ks) by falling head method and some physical and chemical properties:

Ks had highly positive significant correlation with structure coefficient ($R^2 = 0.7409^{**}$), total water stable aggregate > 0.25 mm ($R^2 = 0.6843^{**}$) and positive significant correlation with aggregation index ($R^2 = 0.5726^{*}$), mean weight diameter ($R^2 = 0.5741^{*}$), and highly negative significant correlation with bulk density ($R^2 = 0.7671^{**}$) (Table 4).

Also, multiple regression equation was calculated between the investigated Ks and the five independent variables of physical properties and can be expressed as follows:

 $K_s = 0.375 - 0.324 (B_d) - 7.358 (AI) + 3.676 (MWD) - 0.0305 (S.C) + 0.004 (WSA %) (R² = 0.925[*]).$

The obtained results are confirmed with those found by (Abdel-Nasser, 1995) who found that clay and sand fractions and bulk density are a good estimation of Ks and concluded that mean weight diameter could be used for the predication of the Ks. Also Zein El-Abedine (1979) and Talha et al. (1979) found that the bulk density had a significant and negative effect on the hydraulic conductivity.

	Particular	Linear regression	\mathbf{R}^2	
Y	X	equations		
	WSA %	Y = -0.0755 + 0.0036 X	0.6834	
	Bd	Y = 0.6246 - 0.3623 X	0.7671**	
	MWD	Y = 0.0878 + 0.0596 X	0.5714	
1	AI	Y = 0.0877 + 0.1198 X	0.5726*	
	S.C.	Y = 0.0643 + 0.0457 X	0.7409**	
	Exchangeable Mg ⁺⁺	Y = 0.0219 + 0.0119 X	0.9114	
Ks	Exchangeable Ca ⁺⁺	Y = 0.0288 + 0.0076 X	0.5119	
cm/h	(meq/100 gm soil)			
	(C.E.C.) meq/100 gm soil	Y = -0.0152 + 0.0058X	0.7074**	
	CaCO ₃ %	Y = 0.075 + 0.0582 X	0.8106**	
	ECe (dS/m)	Y = 0.3239 - 0.055 X	0.5935**	
	Soluble Mg^{++} (meq/L)	Y = 0.2195 - 0.0049 X	0.6192**	
	E.S.P	Y = 0.2608 - 0.0123X	0.6804**	
	O.M %	Y = 0.1142 + 0.0489 X	0.7359**	

Table 4. The linear regression between (Ks) determined by falling head method and some physical and chemical properties.

Highly significant at level 1%

* Significant at level 5%.

Ks had highly negative significant correlation with electrical conductivity ($R^{2}=0.5935^{**}$), soluble magnesium ($R^{2}=0.6192^{**}$), exchangeable sodium percentage ($R^{2}=0.6804^{**}$) and highly positive significant correlation with calcium carbonate ($R^{2}=0.6104^{**}$), exchangeable magnesium ($R^{2}=0.9114^{**}$), cation exchangeable capacity ($R^{2}=0.7074^{**}$) and organic matter content ($R^{2}=0.7359^{**}$) and positive correlation with exchangeable calcium ($R^{2}=0.5119^{*}$). (Tables 4).

From data obtained it was noticed that the (Ks) increased with increasing exchangeable Ca, exchangeable Mg, (C.E.C), (CaCO₃%) and organic matter content, while it decreased with increasing electrical conductivity (ECe), soluble Mg, and exchangeable sodium percentage. Therefore, it could be concluded that the previous chemical properties of normal soils are the properties, to which (Ks) is most directly related.

Also, multiple regression equation was calculated between (Ks) of the investigated soils and seven independent variables of chemical properties it expresses as:

Ks = 0.399 + 0.0322 (O.M %) - 0.029 (CaCO₃%) -0.031 (ECe)-8.321 E-05 (soluble Mg⁺⁺) + 0.0267 (exchangeable Mg) -0.007 (C.E.C.)-0.0019(ESP) (R² = 0.999^{+}). Similar results are obtained by (Abdel-Rasoul et al. (1999). They found that the hydraulic conductivity has significant and negative related to ESP. Also, Madkour et al. (1999) showed that Ks has positive significant correlation with organic matter %.

4- The relationship between Ks by auger hole method and by constant and falling head methods:

Simple correlation were calculated to predict the relation between the different three methods for determining hydraulic conductivity, i.e. auger hole method, constant-and falling head method in normal soils. It showed no significant correlation between the auger-hole method and both constant and falling head methods.

REFERENCES

- Abdel-Nasser, G. (1995). Statistical analysis in soil physics studies. 2. Predication of saturated hydraulic conductivity from particle size distribution and soil structure date. J. Agric. Res. Tanta Univ. 21(4): pp. 787-794.
- Abdel-Rasoul, Sh.M.; El-Desouky, H.I. and Hamouda, A.M. (1999). Hydraulic conductivity relations in some soils of Sharkiya Governorate. Egypt, J. Soil Sci. 39 (2): 251-261.
- Arya, L.; Dierolf, T.; Sofyan, A.; Widjaja, A.; IPG; Genuchten-M,k van (1999). Field measurement of the saturated hydraulic conductivity of a macro porous soils with unstable subsoil structure. Soil Sci. 113: (11): 841-852.
- Baver, L.D.; Gardener, W.H. and Gardner, W.R. (1972). Soil physics. John Wiley & Sons, Inc., New York 4th edition.
- Camplell, G.S. and Camplell, M.S. (1982). Irrigation scheduling using soil moisture measurements theory and practice. Advances in Irrigation 1: 25-42.
- Chen, Y.; Tessier, S. and Gallechand, T. (1998). Estimates of tillage effects on saturated hydraulic conductivity Canada. Agric. Eng. Vol. 40, ISS 3; pp. 169-177.
- El-Henawy, A.S. (2000). Impact of available water sources at North Delta on soil and some field crops. M.Sc. Thesis Fac. of Agric. Kafr El-Sheikh, Tanta Univ.
- El-Sammanoudi, I.M. (1992). Impact of applying drainage water on hydrophysical properties of clayey soils in Fayoum depression. Egypt, J. Soil Sci. 32 No. 3 pp. 373-390.
- El-Shafei, Y.Z and Ragal R.L. (1975). Soil surface sealing caused by rain drop impact. Egypt J. soil Sci., 16(1):47-68.

- Gallichand, J.; Madramootoo, C.A.; Enright, P.; Barrington, S.F. (1990). An evaluation of the Guelph permeameter for measuring saturated hydraulic conductivity. Transactions of the ASAE, 33(4): 1179-1184
- Gee, G.W. and Bauder, J.W. (1986). Particle size analysis, In Methods of soil Analysis. C.E. Klute A. (ed.). Part 1. Agron, 9, 15: 383-409. Amer Soc. Agron. Madison, Wisconsin, USA.
- Ibrahim, S.A. (1964). Studies on the size distribution of water stable aggregates in the soils of Nile Delta. M.Sc. Thesis, Ain Shams Univ. Egypt.
- Jackson, M.L. (1958). Soil chemical analysis. Constable Co. Ltd. London.
- James, D.W.; R.J. Hanks and J.H. Jurinak (1982). "Modern irrigated soils". Johnwiley and sons, NY.
- Klute, A. (1986). Methods of soil analysis (part 1). American Society of agronomy, Inc. soil Sci. Soc. of Amer., Inc. Madison, Wisconsin USA 3rd edition.
- Kohnke (1968). Soil physics, Tanta Mc Graw-Hill Pub. Comp. Ltd., New Delhi.
- Madkour, H.E.A.; Hamouda, A.M.M. and Moussa, S.A. (1999). Hydraulic conductivity and structure factor of alluvial soils as affected by some soil parameters. J. Agric. Sci. Mansoura Univ. 24(8): 4249-4258.
- Mostafa, A.T. (1969). The effect of texture, soluble salts and exchangeable cations and the physical properties related to pore-space. Ph.D. Thesis, Fac. Agric., Cairo Univ.
- Oster, J.D.; Shainberg I. and Wood J.D. (1980). Flocculation value and gel structure of Na/Ca montmorillonite and illite suspensions. Soil Sci. Soc. Am. J. 44: 955-959.
- Piper, C.S. (1950). Soil and plant analysis. Inter-Science Publication. New York.
- Richards, L.A. (1954). Diagnosis and improvement of saline and alkali soils. U.S. Salinity Laboratory Staff, Agric. Handbook, No. 60.
- Saffan, M.M. (1984). Zur kenntnis der Boden in der Region Schalma, (Nord Nildelta/Agypten) Ph.D. Thesis, Giessen Uni, Deutschland.
- Saxton, K.E.; Ramls, W.J.; Ranherger, J.S. and R.I. Papendick (1986). Estimating generalized soil water characteristics from texture. Soil Sci. Soc. Am. J. 50: 1031-1036.
- Snedecor, G.W. and Cochran, D.W. (1967). Statistical methods. 6th ed. Oxford and IBH Pub. Co., Calcutta, India.
- Talha, M.; Hamdi, H. and Omar, M.S. (1979). Water movement in calcareous soils under saturated flow. Egypt. J. Soil Sci. 19(1): 73-87.
- Van Beers, W.F.J. (1958). "The auger hole method. A field measurement of the hydraulic conductivity of soil below the water table. Bull. I., Intern, Inst. for Land Reclamation and Improvement. Wageningen, The Nether Lands [C.F. Agric. Res. Rev., Cairo, Vol. 65 No. (4), 63-74].

- Vomocil, J.A. (1957). Measurements of soil bulk density and penetrability. A Review of Methods Adv. Agron. 9: 159-176.
- Yoder, R.E. (1936). A direct method of aggregate analysis of soil and a study of the physical nature erosions losses. J. Amer. Sci. Agron, 28: 337-351.
- Yosry, S.E. (1992). Predication of moisture content at different tension and saturated hydraulic conductivity for some Egyptian soils. M.Sc. Thesis, Faculty of Agriculture, Alexandria University, Egypt.
- Zein El-Abedine, I.A.; Shawky, M.E.; El-Sammanoudy, I.M. (1979). A preliminary study of the effect weight load on some soil properties. 1- Hydraulic conductivity. Res. Bulletin, 1211, Egypt.

الملخص العربي

تأثير الخواص الطبيعية والكيمياتية للأرض العادية على التوصيل الهيدروليكي المقدر بطرق مختلفة

تم أخذ أربعة قطاعات من الأراضي العادية (غير الملحية وغير الصودية) من مركز بيلا محافظة كفرالشييخ ، وقيد تم تتذير بعض الخواص الطبيعية والكيميانية فى هده القطاعات . وقد تم عمل التحليل الإحصبائي بطريقة الانحدار الخطى البيط والمتعدد لجميع القيم المقدرة للتوصيل الهيدروليكي فى الحالة المشيبعة بالطبرق الثلاثة (الضاغط الرأسي الثابت ، الضاغط الرأسي المتغير ، طريقة حفرة الأوجر) ، وجميع الخواص الطبيعية والكيميانية المقدرة فى تلك القطاعات.

ويمكن تلخيص أهم النتائج فيما ياتى:

- 1- التوصيل الهيدروليكي في الحالة المشبعة والمقدر بطريقة الضاغط الرأسي الثابت تأثر بحالة التحبب للتربة حيث وجد ارتباط موجب عالي المعنوية مع كل من المجمعات الثابتة في الماء ذات القطر أكبر من ٢٠, ٥م (**14.88.0 = 2°) و دليل التحبب (**6.03.0 = 2°) و نصف القطر المحوزون (**6.03.0 = 2°). هذا وقد كانت العلاقة معنوية موجبة بين معامل المحوزون (**6.03.0 = 2°). هذا وقد كانت العلاقة معنوية موجبة بين معامل المحوزون (**6.03.0 = 2°). هذا وقد كانت العلاقة معنوية موجبة بين معامل المحوزون (**6.03.0 = 2°). هذا وقد كانت العلاقة معنوية موجبة بين معامل المحوزون (**6.03.0 = 2°). هذا وقد كانت العلاقة معنوية موجبة بين معامل المحيد بين التوصيل الهيدروليكي المشبع والأربعة خواص المستقلة للبناء قد أوضحت عدم وجود المحتعدد بين التوصيل الهيدروليكي المشبع والأربعة خواص المستقلة للبناء قد أوضحت عدم وجود الرتباط معنوي. كذلك فقد وجد أن هناك ارتباط موجب عالي المعنوية بين التوصيل الييدروليكي المشبع والأربعة خواص المستقلة للبناء قد أوضحت عدم وجود المتسبع وكل من الكالسيوم الذائب (**6.03.0 = 8°). والماغنسيوم المتبادل (**7.5.0 هذا المعنوية بين التوصيل الهيدروليكي المشبع والأربعة خواص المستقلة للبناء قد أوضحت عدم وجود المتسبع وكل من الكالسيوم الذائب (**6.03.0 هوجب) والماغنسيوم المتبادل (**7.5.0 هوجه). وقد وجد أي هناك ارتباط موجب عالي المعنوية بين التوصيل الييدروليكي المشبع وكل من الكالسيوم الذائب (**6.03.0 هوجه) والماغنسيوم المتبادل (**7.5.0 هوجه). وعلى المشبع وكل من الكالسيوم الذائب (**6.03.0 هوجه) والماغنسيوم المتبادل (**7.5.0 هوجه). وعلى المشبع وكل من الكالسيوم الذائب (**6.03.0 هوجه) والماغنيوم الذائب (**6.03.0 هوجه) والماغنيوم الكاية (*-1.5.0 هوجه). وعلى المشبع وكل من الكالسيوم الذائب (**6.03.0 هوجه) والماغنوي ماليونية الكالسيوم الذائب (**6.03.0 هوجه) والماغنون والتوصيل المشبع وكل من المعنوي موجه مع كربونات الكالسيوم الكاية (**6.03.0 هوجه). وعلى الهيدروليك والخوام معنوي موجه الماغنوي الخواص الكوميانية الجوسيل الهيدروليك والماغنوي والمعنوي للأدض أن بينهما ارتباط معنوي (**6.09.0 هوجه). وكان الكالسيوم الذائب هم عو أهم الخواص الكيمائية المؤثرة في التوصيل الهيدوميل الهيدروليك ألمي الموبروليكي الأرض أن بينهما ارتباطم الكيميانية ألموشي والوكي الموسيل اله
- ٢- للتوصيل الهيدروليكي المشبع والمقدر بطريقة الضاغط الرأسي المتغير فقد تأثر بشدة بحالة التحبب فــي الأرض حيث وجد التوصيل الماني يرتبط ارتباط موجب عالي المعنوية مع كل من المجمعات الثابتة فى الماء والتي أقطارها لقل من ٢٥,٠٥م (**R2=0.6824) وكذا معامل البناء (**R2=0.7409) وقد لوحظ أن هناك ارتباط معنوي موجب مع كل من نصف القطر الموزون (*R2=0.5714) ودليل البناء (*R2=0.5726=2).

وبالنسبة للكثافة الظاهرية فقد أظهرت النتائج أنها تؤثر بدرجة كبيرة على التوصيل الهيدروليكي لمشبع حيث هذك لرتبلط سلب على لمعنوية بين التوصيل الماتي والكثافة لظاهرية للأرض (**80.761). وقد أوضـحت معادلة الانحـدار الخطـي المتعدد بين الخواص الطبيعية المذكورة سابقا والتوصيل أوضـحت معادلة الانحـدار الخطـي المتعدد بين الخواص الطبيعية المذكورة سابقا والتوصيل الهيدروليك. أن هـناك ارتـبلط معـنوي (*80.929) بينهما. وكان من الواضح أن الكثافة الظاهرية هـي أهـم العوامـل التي تؤثر في التوصيل الهيدروليك. وقد كان واضحا أن الخواص الطبيعية المذكورة سابقا والتوصيل الهيدروليكي أن هـناك ارتـبلط معـنوي (*80.929) بينهما. وكان من الواضح أن الكثافة الظاهـرية هـي أهـم العوامـل التي تؤثر في التوصيل الهيدروليكي. وقد كان واضحا أن الخواص الرأسي المتغير حيث كان هناك ارتـبلط موجب على التوصيل الهيدروليكي المقدر بطريقة الضاغط الرأسي المتغير حيث كان هناك ارتباط موجب على المعنوية بين التوصيل الهيدروليكي المقدر بطريقة الضاغط الرأسي المتغير حيث كان هناك ارتباط موجب عالي المعنوية بين التوصيل الهيدروليكي المقدر بطريقة الضاغط الرأسي الماني المتي حيث الموسيل المالي الهيدروليكي المقدر بطريقة الضاغط الرأسي المتغير حيث كان هناك ارتباط موجب عالي المعنوية بين التوصيل الهيدروليكي وكل من الكاسيوم المتبادل (**10.94) والسعة التبادية الكاتيونية (**10.709) ، كربونات الرأسي الماغلية(**10.800) والمعة التبادية الكانيونية (**10.800) ، كربونات أن الكاسيوم المتبادل كان له تأثير موجب على التوصيل المتي التربة (**10.800) وحتوى المادة العضوية للأرض (**10.800) والسعة المتوي المتربة (لمتربة (**10.800) وجلامى لمتربة المتوية المانين مالينوي ماليان ليزيادة أن الكاسيوم المتبلال كان له تأثير موجب على التوصيل المتي التربة (**10.800) وحتوى المادة العضوية المارض (لماد الخواص الخواص الحلومي المور الكاليوم المتبلال كان لويادة أن الكاسيوم المتبلال كان له تأثير موجب على التوصيل المتي التربة (**10.800) وحد كان ليزيادة أن الخواص الكاسيوم المتبلال كان لواليوم (**10.800) وحدوى المتوم المتبادل تأثير معنوي سالب على التوصيل الميدروليكي للأرض (**10.800) وحالي المتبادل تأثير معنوي سالم على التوصيل الميومي المياليومي المومي المان لينا مانوميكي الأرض (**10.800) وحدوى المومي الماني واليوم واليكي والي

٣- كذلك دلست النستانج على عدم وجود ارتباط معنوي بين قيم التوصيل الهيدروليكي المتحصل عليها بطريقة الضاغط الرأسي للثابت والضاغظ الرأسي المتغير أو بطريقة حفرة الأوجر.