STUDIES ON SOILS OF TOSHKA REGION 2- FORMS OF INORGANIC PHOSPHATE IN SOILS OF TOSHKA REGION

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ABSTRACT: Fifteen soil profiles were dug, during the 2001 year, at Toshka region and forty five soil samples were collected from layers 0-30, 30-60, and 60-90 cm.

The inorganic P-fractions decreased with soil depth increases down the profile, and varied greatly along with the soil depth.

The relative abundance of P-fractions in most soils is in the order: Ca-P> saliod-P \approx Fe-P > Al-P. The Ca-P was the most dominant fraction in these soil.

Fine textured soils contained more total – and saloid with the coarse ones. High positives correlations were obtained between the clay content and either total or available phosphorus.

INTRODUCTION

Toshka region is bounded by longitudes 30° 30' and 32° 00' east., and latitudes 22° 22' north (Said, 1962). In this region work is under way, to extend the Nile water into the heart of the western desert to the Dakhela and Kharga oases. This bold project would create in effect a new agricultural land and was described as the national project for the twenty first century, (Egyptian Ministry of Economy, Dec. 1998).

Generally, desertlands are sandy in texture, poor in both organic matter and available nutrients content (Mitkees *et al.*, 1966 a,b,c; Abdel Latif *et al.*, 1987; Khatab *et al.*, 1999, Rahmou *et al.*, 2001and Hassanein *et al.*, 2002).

Phosphorus in soils can be classed, generally as organic and inorganic, depending on the nature Behiry, et. al.

of the compounds in which its occurs. The organic fraction is found in humus and other organic materials, (Black, 1968). The inorganic fraction occurs in numerous combinations with iron, aluminum, calcium, fluorine and other elements. These compounds, are usually only very slightly soluble in water as reported by (Chang and Jackson, 1957; Behiry, 1991 and Diab *et al.* 1994). Phosphorus, also reacts with clay to form generally insoluble clay-phosphate complexes (Tahoun *et al.*, 1975).

According to (Cooke, 1970), total soil phosphorus is often a useful parameter of soil-phosphate capacity. Soil poor in total phosphorus, usually has little capacity to supply the element to growing plants. Amer and Abou El-Roos, (1975); Ibrahim *et al.* (1980) and Diab *et al.* (1994) reported that total phosphorus in the Egyptian soils varies considerably, the low values are associated with sandy and calcareous soils and the high values are associated with heavy-textured alluvial soils.

The aim of this work is to study the individual fractions of phosphorus content in the virgin soils of Toshka region.

MATERIALS AND METHODS

Fifteen soil profiles were dug, during the year 2001, at Toshka region and soil samples were collected from tree layers: surface (0-30 cm), sub surface (30-60 cm) and deep (60-90 cm) layers. The location of the selected area are given in Fig. (1).

Samples were air dried, ground by wooden pestle., and sieved through a 2 mm nylon sieve. Particle size distribution was determined using the method (Kilmer pipette and Alexander, 1949), using sodium hexameta-phosphate as a dispersing agent. Soil pH, in soil paste extract was measured by a glass electrode pH-meter (Richards, 1954). Soil salinity in of electric terms conductivity (Ec) was determined, in the saturation extract. The organic matter content was determined according to Walkely and Black method, as given by Jackson (1958). Calcium carbonate was determined volumetrically suing the calcimeter as described by Piper (1950). Gypsum content also was. determined by precipitation with acetone (Richards, 1954). Available phosphorus was determined by the method of Olsen et al. (1954). Fractionation of soil phosphorus was performed according to the modified method of Chang and Jackson (1957) as follows:

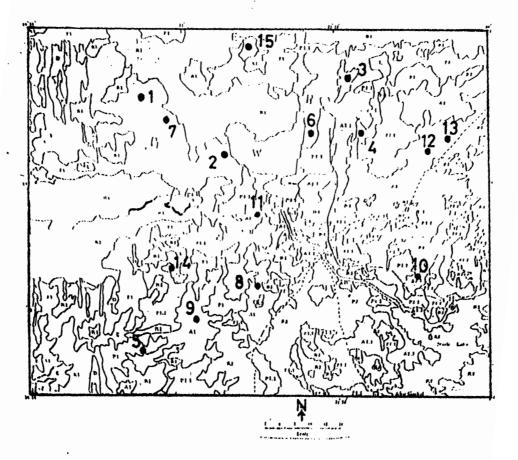


Fig. (1): Locations of the studied soil samples.

- Physiographic soil legend:

- $P_1 \rightarrow Pediplain of soft rock (shale)$
- $A_1 \rightarrow Alluvial deposits over soft rock suprofite$
- $A_2 \rightarrow$ Alluvial deposits over hard rock (sandstone) suprofite
- $L \rightarrow Playa$
- $E \rightarrow Acolian plain$

- Miscelianeous land types and cover:

- $\mathbf{R}_1 \rightarrow \mathbf{Dissected}$ questa.
- $R_2 \rightarrow Rock$ land.
- $\mathbf{D} \rightarrow \text{Barchan sand duncs}$
- $W \rightarrow$ Seeped water body with changeable beach line.

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1- Saloid phosphorus was extracted by 1N NH₄Cl using a soil solution ratio of 1:50.

- 2- Al-phosphorus was extracted using 0.5 N NH₄F at pH 8.5 suing a soil solution ratio of 1:100.
- 3- Fe-phosphorus was extracted using 1.0 N NaOH + NaCl using a soil solution ratio 1:50.
- 4- Ca-phosphorus was extracted by 0.5 N H₂SO₄ using a soil solution ratio of 1:50.

In all phosphorus fractions, the ascorbic acid blue method as given by John (1970) was utilized, to measure P in solution.

RESULTS AND DISCUSSION

Soil properties: Properties of the studied soils are listed in Tables 1 and 2. The mechanical analysis showed that these soils are general. characterized in by medium in texture with a various range of both coarse sand + Fine ranged from (15.2sand % 88.25%). (20.5 - 88.09%)and (26.72 - 87.37%) for the surface subsurface and deep soil samples, respectively. The same pattern was found for both silt + clay % which ranged from (11.75 - 84.8%), (11.91 - 79.50%) and (12.63 -

73.28 %) in the aforesaid soil samples, in a respective order. In general, the pH values were alkaline in the soil samples under study. Data also, showed that these soils are very poor in organic matter content are mainly due to the dryness and hot climate (Moustafa, et al., 1988).

Concerning calcium carbonate, the obtained data reveal that its values widely varied within the studied soil profiles. The highest values of CaCO₃ % were found in soil profiles No: 3, 4, 5, 6, 8, 10 and 11. It could be stated that these seven sites have to be considered as calcareous soils, according to the definition of Anter et al. (1973).

Gypsum content was low, except for two soil sites which had a moderate content i.e., profiles No. 3 and 6. The values of electric conductivity (Ec) were in the range of 1.0 to 48.0 dSm⁻¹ in the surface soil samples, and from 1.0 to 86.6 dSm⁻¹, in the subsurface soil samples, while the deep ones ranged from 2.92 to 35.0 dSm⁻¹, indicating that soil salinity of the studied profiles flucetated between non-saline to extremely saline. In general, the soluble cations were found to be in the order of $Na^+ >$ $Ca^{++} > Mg^{++} > K^{+}$ in the

	Denth	Partic	le size		oution	Silt +	О.М.	CaCO ₃	Gypsum	Texture
Prof.	Depth. Cm,		(%	<u> </u>		Clay			0/	
No.	Cm,	C .	F .	Silt	Clay	%	%	%	%	class
		Sand	Sand	1						
1	0-30	8.0	25.5	27,8	38.8	66.6	0.1	9.8	0.6	CL.
	30-60	1.5	29.6	32.2	36.8	69.0	0.3	6.6	0.3	CL.
	60-90	2.9	25.9	30.8	40.5	71.3	0.2	7.5	0.2	C.
2	0-30	4.2	34.8	16.0	45.1	61.0	0.2	6.5	2.4	C.
	30-60	3.1	30.9	25.2		66.0	0.2	4.1	1.2	C.
	60-90	2.8	37.3	17.8	42.2	60.0	0.2	2.6	1.2	C.
					`	2			• .	
.3	0-30	25.7	34.3	17.2		40.0	0.1	15.2	4.0	SCL.
	30-60	5.1	35.9		30.9	59.0	0.5	23.5	9.7	CL.
	60-90	4.1	32.2	30.5	33.3	63.8	0.4	28.2	3.7	CL.
4	0-30	2.9	12.4	34.2	50.6	84.8	0.1	13.1	1.7	С.
	30-60	4.6	15.9	35.5	44.0	79.5	0.2	13.7	1.5	С.
	60-90	7.1	19.6	27.2	46.1	73.3	0.2	13.9	1.5	С.
5	0-30	15.7	30.0	22.6	31.7	54.2	0.2	14.6	3.0	SCL.
	30-60	21.2	28.9	22.3	27.6	49.9	0.3	15.3	2.0	SCL.
	60-90	28.0	27.8	19.1	25.2	44.3	0.3	19.3	1.7	SCL.
6	0-30	52.9	31.0	7.4	8.7	16.1	0.1	13.7	9.4	SCL.
·	30-60	50.9	27.8	8.1	13.3	21.4	0.2	12.5	4.2	SL.
	60-90	14.5	24.6	24.3			0.1	10.1	2.7	CL.
7	0-30	29.1	28.3	18.1	24.6	42.6	0.1	8.1	1.5	SCL.
	30-60	16.3	25.6	25.1	33.2	58.2	0.2	4.4	0.9	CL.
	60-90	18.2	27.3	19.3	35.3	54.6	0.2	5.0	0.5	SC.
8	0-30	41.0	28.8	12.9	17.3	30.2	0.1	21.2	0.4	SL.
	30-60	54.3	22.1	10.6	13.1	23.7	0.3	10.8	0.3	SL.
	60-90	50.6	28.7	11.2	9.6	20.8	0.1	10.8	0.2	SL.

Table (1): Particle size distribution, organic matter (O.M.), CaCO₃ and gypsum content in the studied soil profiles.

Table (1): Cont.

Prof.	Depth.	Partic	le size (%		bution	Silt + Clay	O.M .	CaCO ₃	Gypsum	Texture
No.	Cm,	C. Sand	F. Sand	Silt	Clay	%	%	%	%	class
9	0-30	46.7	2 7.9	10.7	14.8	25.5	0.1	1.2	0.3	SL.
	30-60	50.3	26.2	9.9	13.6	23.5	0.1	1.5	0.2	SL.
	60-90	46.7	30.3	10.3	12.7	23.0	0.1	2.8	0.3	SL.
10	0-30	45.3	23.4	12.7	18.9	31.3	0.1	20.0	0.4	SL.
	30-60	49.7	22.1	11.7	16.5	28.2	0.1	22.0	1.3	SL.
	60-90	45.5	25.8	12.0	16.7	28.7	0.1	24.6	4.2	SL.
11	0-30	66.1	22.2	5.1	6.7	11.8	0.0	14.9	0.3	S.
	30-60	61.2	25.0	6.1	7.7	13.8	0.0	10.7	0.3	S.
	60-90	43.6	28.9	12.3	15.2	27.5	0.0	12.3	0.3	SL.
12	0-30	59.6	27.2	5.9	7.4	13.3	0.0	10.8	0.3	LS.
	30-60	56.2	31.2	5.5	7.1	12.7	0.0	5.9	0.2	S.
	60-90	51.9	24.4	9.5	14.3	23.8	0.1	3.6	0.2	SL.
13	0-30	62.8	25.4	4.8	6.9	11.8	0.0	9.1	0.3	S.
	30-60	54.6	32.1	5.8	7.9	13.6	0.0	6.2	0.3	LS.
	60-90	50.1	33.0	6.8	10.1	16.9	0.1	9.4	0.3	LS.
14	0-30	65.0	20.4	7.2	7.4	14.6	0.1	12.1	0.4	LS.
	30-60	67.8	17.4	6.7	8.1	14.8	0.1	10.2	0.3	LS.
	60-90	62.2	20.9	7.1	8.7	15.8	0.1	10.6	0.3	LS.
15	0-30	55.3	31.9	5.6	7.2	12.8	0.0	3.7	0.4	S.
	30-60	59.6	28.5	5.1	6.8	11.9	0.1	1.3	0.6	S.
	60-90	56.6	30.8	5.4	7.2	12.6	0.1	0.4	0.4	S.

C: Clay; LS: Loam sand;

CL: Clay loam; SCL: Sandy clay loam

SL: Sandy loam

Prof.	Depth.	pН	EC	s	Soluble ions in saturation extract (meq/L)						
No.	Cm,		dSm ⁻¹		Catio	the second s		Anions			mg Pkg ⁻¹
				Ca ⁺⁺	Mg ⁺⁺		\mathbf{K}^{+}	HCO ₃	and the second se	SO4	soil
1	0-30	7.4	31.8	131.1	44.7	250.0	2.8	1.8	264.0	162.7	17.1
	30-60	7.7	13.3	31.6	23.2	122.5	2.1	2.2	89.0	88.2	15.4
	60-90	7.3	35.0	121.1	60.0	280.0	3.3	2.0	299.0	163.3	12.0
2	0-30	7.6	8.2	32.0	5.9	66.2	0.8	2.0	47.6	55.2	18.9
	30-60	7.6	6.5	46.1	13.9	24.6	0.8	2.1	32.9	50.3	17.2
	60-90	7.5	10.0	41.3	7.0	88.0	1.0	2.0	58. 7	76.5	17.7
3	0-30	7.6	8.5	27.4	5.0	72.5	2.1	2.4	19.0	85.5	13.6
	30-60	8.0	22.1	35.3	10.0	242.5	3.4	1.6	107.0	182.5	13.0
	60-90	7.8	17.2	41.1	16.1	172.5	2.2	1.6	100.0	130.2	10.0
4	0-30	7.9	7.4	25.3	2.8	67.5	1.1	1.8	7.0	87.9	21.2
	30-60	7.8	19.4	32.6	9.8	185.0	1.5	1.6	135.0	92.2	19.5
	60-90	7.7	15.4	36.8	7.0	137.5	1.4	1.6	111.0	70.1	18.4
5	0-30	7.5	10.6	36.3	11.3	88.8	1.7	2.2	44.0	91.8	13.3
	30-60	7.5	8.9	30.0	8.6	70.0	1.3	2.0	35.0	72.9	11.6
	60-90	7.5	7.8	30.0	7.1	63.8	1.4	2.2	25.0	75.1	10.6
6	0-30	7.2	48.0	233.7	25.4	390.0	2.6	2.0	400.0	249.7	3.7
	30-60	7.2	86.6	276.2	43.3	845.0	1.1	2.0	1098.0	95.6	5.6 ·
	60-90	7.6	26.1	41.8	20.8	292.5	1.7	1.3	205.0	150.5	5.4
7	0-30	7.6	10.0	35.3	17.6	67.5	1.9	1.6	49.0	71.6	14.3
	30-60	7.6	17.2	31.6	27.5	147.5	1.4	1.8	106.0	100.1	13.9
	60-90	7.4	17.0	47.4	32.6	125.0	1.3	1.4	128.0	76.9	10.8
8	0-30	7.8	3.2	16.3	4.6	20.0	1.7	1.6	7.0	34.1	7.3
	30-60	8.0	5.6	9.0	2.0	53.8	1.4	1.6	19.0	45.5	5.5
	60-90		14.8	40.5		135.0	1.8	1.4	93.0	94.7	4.0

Table (2): Chemical composition of the soil saturation extract of the studied soil profiles.

Table (2): Cont.

	Depth.	pН	EC	Soluble ions in saturation extract (meq/L)							Available Phosphorus
No.	Cm,	e	dSm ⁻¹		Cati	_		_	Anions	mg Pkg'	
.				Ca ⁺⁺	Mg ⁺⁺	Na ⁺	\mathbf{K}^{+}	HCO ₃	Cľ	SO ₄ ⁼	soil
9	0-30	7.7	5.1	29.0	2.3	34.0	0.9	1.7	14.5	49.9	6.2
	30-60	7:4	5.1	36.8	1.3	25.0	0.9	1.2	22.0	40.8	5.7
	60-90	7.4	5.2	37.9	5.4	19.3	0.8	1.2	26.0	36.2	5.4
10	0-30	7.7	15.7	41.1	9,9	130.0	3.6	1.6	89.0	93.9	7.8
	30-60		20.3	74.2		150.0		1.8		117.3	6.9
	60-90	7.6	22.8	59.6		195.0		1.8		134.5	7.0
11	0-30		1.0	3.7	0.9		0.8	2.6	1.7	5.9	6.7
	30-60		2.5	6.3	1.5	17.1		2.0	10.5	13.5	3.2
	60-90	7.4	5.8	32.5	9.9	16.6	1.1	2.4	30.8	26.9	2.4
12	0-30	8.1	1.4	4.7	1.0	9.2	0.8	2.3	3.5	9.9	5.1
	30-60	8.3	1.0	3.2	1.1	9.2	0.4	2.0	2.6	9.3	3.0
	60-90	7.4	6.1	30.5	6.6	32.0	1.2	1.4	28.0	41.0	4.0
13	0-30	8.1	1.2	7.7	1.1	5.2	0.7	2.0	2.2	10.4	4.9
	30-60		6.5	36.8	3.6	30.0		1.4	26.0	44.6	3.3
	60-90	7.6	2.9	15.3	1.4	14.1		1.4	13.5	16.5	3.1
14	0-30		6.8	26.3	9.9	53.8		1.8	35.0	55.1	4.0
	30-60		12.9	50.0	11.9	86.3		1.6	7 2.0	76.3	3.4
	60-90	7.8	8.3	40.0	7.4	<u>54.0</u>	1.4	1.4	40.0	61.1	3.7
15	0-30	7.8	4.7	31.6	4.6	14.0	3.2	1.4	3.0	49.0	3.0
	30-60	7.6	11.1	43.7	10.1	112.5	3.0	1.6	61.0	106.7	2.9
	60-90	7.3	15.8	95.5	29.5	56.0	1.7	1.2	123.0	58.5	2.0

investigated soil samples. In addition, the soluble anions showed a general distribution pattern of Cl- > SO₄= > HCO₃, while CO₃⁻ was absent in the soil samples. Similar results were found in the 1st, part of the current study about Toskha region (Hassanein, *et al.*, 2002).

Inorganic phosphorus fractions: The obtained results of the inorganic phosphorus fractions in the investigated soils are given in Table (3). These data indicate that all phosphorus fractions in all soils gradually decreased as the soil depth increased down in the profile. These results agree with those obtained by Piccolo and Huluka (1986); Fuleky, (1997) and Abdel Hamed (1998). This effect is attributed to the low mobility of soluble phosphorus in most soils. Sorensen and Wiese (1982) and Behiry (1991) reported that the upper layers of soils usually contain more organic and inorganic phosphorus than the lower layers.

These data indicate that the amount of individual phosphorus fractions varies greatly along with the soil depth. The relative abundance of these fractions in most soils is the order: Ca-P > Saloid-P \approx Fe-P > Al-P.

Table 3 also shows that the total inorganic phosphorus, (including Saloid-P, Al-P, Fe-P and Ca-P,)

varied greatly in the soils with profile means ranging from 41.5 mg Pkg⁻¹ soil, in the sandy soil (profile 15) to 844.2 mg Pkg⁻¹ in the clay soil (Profile, 3) Phosphorus was highest in the clayey soils than the sandy ones, while the loamy soils were in between. Pierzynski et al. (1990) reported that the total phosphorus concentration is. generally, higher in the clay fraction of the soils compared with the silt or the sand fractions. Abdel Hamed (1998) found that means of total inorganic phosphorus in profiles ranged from 321.0 to 661.4 mg/kg soil in Dakhlia alluvial soils. The obtained values of inorganic phosphorus are in well agreement with those obtained by several Egyptian workers such as Mashady and Ramy (1981), Diab et al. (1994) and Abdel Hamed (1998).

As regards the individual forms of phosphorus in the soil. the aforementioned table, shows that Ca-P content is far greater than any other fraction, and this may be due to; a) the alkaline pH reaction; b) the high content of calcium carbonate in these soils (specially soils of profiles: 3, 4, 5, 6, 8, 10, 11 and 14); and c) the calcium ions are often high in the soil solution of such soils. Chang and Jackson (1957),

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Prof.	Depth.	Saloid-P	Al-P	Fe-P	Ca-P	Total
No.	Cm,		(mg	Pkg ⁻¹)		iual
1	0-30	7.1	4.7	2.2	558.8	572.8
	30-60	6.7	0.6	1.6	552.5	561.4
	60-90	5.0	0.3	0.9	537.5	543.7
	Mean	6.3	1.9	1.6	549.6	559.4
2	0-30	6.9	4.9	17.1	920.0	948.9
	30-60	6.1	0.7	13.7	750.0	770.9
	60-90	4.8	0.5	11.3	575.0	770.5
	Mean	5.9	2.0	14.0	748.3	770.2
3	0-30	6.3	4.1	4.1	777.5	792.0
	30-60	5.3	0.6	1.3	597.5	604.7
	60-90	4.1	0.3	0.9	507.5	512.8
	Mean	5.2	1.7	2.1	627.5	636.5
4	0-30	7.9	2.2	12.5	925. 7	948.3
	30-60	6.7	1.9	9.8	735.0	753.4
	60-90	3.0	0.7	7.1	820.2	831.0
•	Mean	5.9	1.6	9.8	826.9	844.2
5	0-30	4.8	2.7	13.9	623.8	645.0
	30-60	4.5	0.9	5.2	615.0	625.6
	60-90	4.0	0.2	4.2	475.0	483.4
<u> </u>	Mean	4.3	1.3	7.8	571.3	584.7
6	0-30	8.3	3.2	3.1	580.0	594.6
	30-60	7.6	0.5	1.8	475.0	484.9
	60-90	4.8	0.5	1.8	167.8	174.9
	Mean	6.9	1.5	2.0	407.6	418.0
7	0-30	10.4	0.8	11.5	623.8	646.5
	30-60	9.6	0.7	3.8	560.3	574
	60-90	0.1	0.5	3.5	115.8	119.9
	Mean	6.7	0.7	6.2	435.3	452.9
8	0-30	7.9	0.5	5.3	435.0	448.7
	30-60	6.0	0.3	- 3.8	95.0	105.1
	60-90	3.1	0.2	1.6	91.4	96.3
	Mean	5.6	0.4	3.6	207.3	216.9

 Table (3): Inorganic phosphorus fractions (mg Pkg⁻¹ soil) in Toshka soil samples.

Prof.	Depth.	Saloid-P	Al-P	Fe-P Pkg ⁻¹)	Ca-P	Total
No.	Cm,		I Otal			
9	0-30	7.3	2.2	4.5	565.0	579.0
	30-60	5.9	0.9	3.5	559.5	569.8
	60-90	5.1	0.7	2.0	544.6	552.4
	Mean	6.1	1.3	3.3	556.4	567.1
10	0-30	6.0	3.5	1.5	615.0	626.0
	30-60	5.1	0.3	0.8	560.0	566.2
	60-90	4.3	0.3	0.8	552.5	557.9
	Mean	5.1	1.4	1.0	575.8	583.3
11	0-30	9.3	3.0	15.3	200.6	228.2
	30-60	8.2	2.0	11.2	79.0	100.4
	60-90	7.9	0.3	9.9	66.0	84.1
	Mean	8.5	1.8	13.1	115.4	138.8
12	0-30	7.7	2.3	14.5	244.4	268.9
	30-60	6.5	0.6	11.1	107.5	125.7
	60-90	3.6	0.5	9.8	60.3	74.2
	Mean	5.9	1.1	11.8	137.4	156.2
13	0-30	8.4	3.5	12.0	177.3	201.2
	30-60	7.0	1.3	10.5	151.1	169.9
	60-90	7.9	0.2	9.1	129.8	147.0
	Mean	7.8	1.7	10.5	152.7	172.7
14	0-30	8.9	2.2	16.1	260.1	287.3
	30-60	6.3	1.6	7.0	180.2	195.1
	60-90	6.2	0.8	2.5	100.0	109.5
	Mean	7.1	1.5	8.5	180.1	196.3
15	0-30	6.8	2.1	10.0	28.0	46.9
	30-60	5.5	0.9	8.2	27.0	41.6
	60-90	3.6	0.7	5.5	25.3	41.4
	Mean	5.3	1.2	7.9	27.3	41.5
Grand		6.17	6.2	1.4	6.9	366.5

Table (3): Cont.

Behiry, el. di

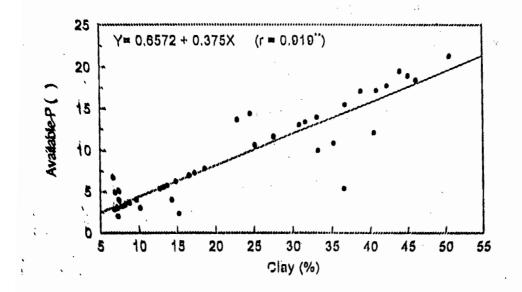
Lindsay and Moreno (1960) and Behiry (1991) reported that the Ca-P fractions are more likely to be formed during soil weathering, as well as by fixation and precipitation of applied soluble phosphate fertilizers. The obtained data clearly shows that the Ca-P fraction is the most dominant in the investigated soils. It ranged from 27.1 to 826.9 mg Pkg⁻¹ soil. The Ca-P represent 65.3 to 97.9% of the total inorganic phosphorus. Diab et al. (1994), found that more 60% of phosphates exists in the form of Ca-P, while, Abdel Hamed (1998) found that 80% of the total inorganic phosphorus was in that form.

The former data, also reveal that of Fe-P which represent fine crystals of natural iron phosphate minerals, such as stregite and dufenite or precipitated phosphate on iron oxides surface (Chang and Jackson 1957), is very small compared with Ca-P fraction. The obtained mean values of Fe-P ranged between 1.0 and 14.0 mg Pkg⁻¹ soil, with a grand mean of 1.4 mg Pkg⁻¹ soil. This indicates that the formation of Fe-phosphate is unfavoured under alkaline soil conditions especially with presence of little amounts of free iron oxides as reported by Sing et al. (1966). Hola et al. (1985) reported that in soils of recent,

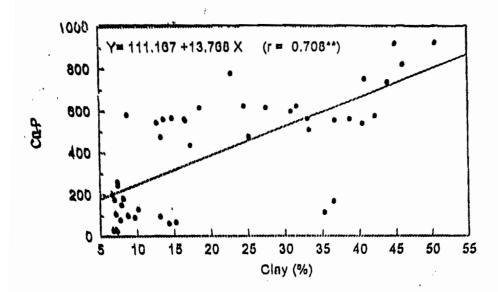
origin calcium and aluminum phosphate are more likely to be found than iron phosphate. This is attributed to the relatively higher activities in the soil of Ca-and Al-ions than Fe-ions, which are controlled by the activities of the respective cations of CaCO₃, alumino-silicates, gibbsite and iron oxides.

Table (3) shows that the mean values of saloid-P in soil profiles ranged between 4.3 to 8.5 mg Pkg⁻¹ soil, in the investigated soil samples. These results are similar to those found by El-Rashidy (1972), Amer and Abou Roos (1975) and Abdel Hamed (1998). This fraction does not represent a certain soluble state and was defined by Hesse (1971) as the soluble and loosely bound –P. which seems to represent the immediately available phosphorus fraction in soil.

Al-P which includes phosphorus derived from the Al-P crystals of certain minerals such as variscites, tarankite and wavellite or from aluminum phosphate precipitation on alumino-silicates or gibbsites (Chang and Jackson 1957), wa very small than other Phosphat fractions. Its mean content in th tested profiles ranged from 0.4 1 2.0 mg Pkg⁻¹ soil. Zagazig J.Agric. Res., Vol. 30 No.(5) 2003



rig. (2): The relation between clay content and Ca-phosphorus.



'ig. (3): The relation between clay content and Ca-phosphours.

Available phosphorus. Data in Table (2) showed the available phosphorus content of the investigated soils as determined by the NaHCO₃ method (Olsen' et al., 1954), ranged between 3.0 - 21.2, 2.9-19.5 and 2.5-18.9 mg Pkg⁻¹ soil in the surface, subsurface and deep soil samples, respectively. Generally, fine textured soils contained more than coarse textured soils. Fig. (2) represents the relationship between available phosphorus and clay content. The calculation showed a highly positive correlation between available phosphorus and clay content. These results are similar to those reported by Dahdouh (1981), Pierzynski et al. (1990) and Behiry (1991).

It is interesting to note, that a highly positive correlation coefficient was found between the clay content and Ca-P (Fig. 3). No significant correlation was found between available phosphorus and either CaCO₃ % or organic matter %.

Data of Table 3 shows that the values of available phosphorus of these virgin soils, were low, excepted soils of profiles 1,2 and 4, which contained adequate amounts of available phosphorus according to the fertility interpretion by Soltanpour and Schwab (1977) and Zanati *et al.* (1982).

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٢- صور الفوسفور المعنى في أراضي من منطقة توشكي

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جمعت ٤٤ عينة تربه سطحية وتحت سطحية وعميقة مسن الأسلق ٥-٣٠، ٣٠-٢٠، ٢٠-

٩٠ سم من خمسة عشر قطاع تربه فى منطقة توشكى. وقد قدرت الصفات الطبعية والكيمائية لهذه العينات وأشارت النتائج الى تناقص صور القوسفور المعدى تدريجياً مسع العمق، كما تنوعت كثيراً معه ولقد توزعت تلك الصور كآلاتى:

 $Ca-P > Saloid-P \approx Fe-P > Al-P$

وكان المحتوى السائد هو Ca-P ، كما أظهرت النتائج ان الأراضى الناعمة القوام ذات محتوى عال من القوسفور الكلى والميسر بالمقارنة بالأراضى الخشنة لقوام ، ولقد ارتسبط محتوى الطين إحصانيا معنويا. مع كل من Ca-P، والفسفور الميسر.